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Dynamics of Energy Balance Models for Planetary Climate

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April 13, 2016

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Background • ○ ○ Dynamics of the Ice Line

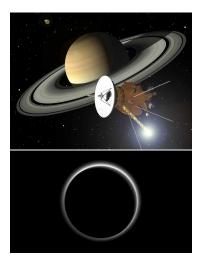
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Motivation

- Low dimensional climate models are important for understanding the predominant forces affecting the climate of Earth
- Fly-bys of Pluto and other rocky celestial bodies in our solar system have raised interest in other climates



Photos from nasa.gov

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Energy Balance Models in 1969

- Budyko and Sellers (independently) proposed energy balance models for the Earth (1, 14)
- Wanted to study if another glacial age was possible
- Both models had the same major components: incoming solar radiation, outgoing radiation, and energy transfer:



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$$R\Delta T = Q(y)(1 - \alpha(y)) - (A + BT) + \Gamma(T)$$

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Energy Balance Models Today

$$R\frac{\partial T}{\partial t} = Qs(y)(1 - \alpha(y, \eta)) - (A + BT) + \Gamma(T)$$

where

$$\Gamma(T(y,\eta)) = -C\left(T(y,\eta) - \int_0^1 T(\gamma,\eta)d\gamma\right).$$

Widiasih introduced an equation for the dynamics of the ice line, η , in 2012 (18)

$$\frac{d}{dt}\eta = \epsilon(T(y,\eta) - T_c)$$

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"Quadratic Approximation"

In (9) McGehee and Widiasih consider the Budyko-type equation for $y,\eta\in[0,1]$

$$\frac{\partial}{\partial t}T = \frac{1}{R}\left(Qs(y)(1-\alpha(\eta,y)) - (A+BT(y,\eta)) - C\left(T(y,\eta) - \overline{T}\right)\right)$$

with dynamic ice line

$$\dot{\eta} = \rho(T(y,\eta) - T_c)$$

and piecewise constant albedo function

$$\alpha(y,\eta) = \begin{cases} \alpha_w & y < \eta \\ \alpha_0 & y = \eta \\ \alpha_i & y > \eta \end{cases}$$

where

$$\alpha_0 = \frac{\alpha_w + \alpha_i}{2}.$$

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Piecewise Function Space

The equilibrium temperature profile can be written

$$T_{\eta}^{*}(y) = \begin{cases} \frac{1}{B+C}(Qs(y)(1-\alpha_{w}) - A + C\overline{T_{\eta}^{*}}) & y < \eta\\ \frac{1}{B+C}(Qs(\eta)(1-\alpha_{0}) - A + C\overline{T_{\eta}^{*}}) & y = \eta\\ \frac{1}{B+C}(Qs(y)(1-\alpha_{i}) - A + C\overline{T_{\eta}^{*}}) & y > \eta \end{cases}$$

which motivates the four-dimensional function space X whose elements are of the form

$$T(y) = \begin{cases} w_0 + \frac{1}{2}z_0 + (w_2 + \frac{1}{2}z_2) p_2(y) & y < \eta \\ w_0 + w_2 p_2(\eta) & y = \eta \\ w_0 - \frac{1}{2}z_0 + (w_2 - \frac{1}{2}z_2) p_2(y) & y > \eta \end{cases}$$

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New Budyko's "Equation"

Reformulating the $\partial_t T$ equation in this function space gives

$$\begin{aligned} R\dot{w}_{0} &= Q(1-\alpha_{0}) - A - Bw_{0} + C\left((\eta - \frac{1}{2})z_{0} + z_{2}\int_{0}^{\eta}p_{2}(y)dy\right) \\ R\dot{z}_{0} &= Q(\alpha_{i} - \alpha_{w}) - (B + C)z_{0} \\ R\dot{w}_{2} &= Qs_{2}(1-\alpha_{0}) - (B + C)w_{2} \\ R\dot{z}_{2} &= Qs_{2}(\alpha_{i} - \alpha_{w}) - (B + C)z_{2} \end{aligned}$$

and the ice line equation becomes

$$R\dot{\eta} = \rho \left(w_0 - \frac{Q(1-\alpha_0)}{B+C} s_2 p_2(\eta) + T_c \right).$$

Write

$$R\dot{w}_{0} = -B(w_{0} - F(\eta))$$

$$R\dot{\eta} = -\rho(w_{0} - G(\eta))$$

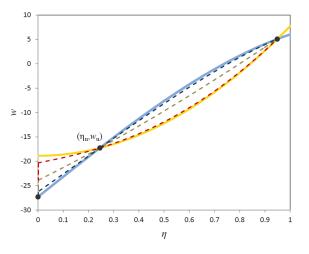
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(η, w_0) Phase Space



$$\begin{aligned} R\dot{w}_0 &= -B(w_0 - F(\eta)) \\ R\dot{\eta} &= -\rho(w_0 - G(\eta)) \\ R\dot{\eta} &= -\rho(w_0 - G(\eta)) \end{aligned}$$

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The Jormungand Model

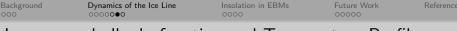
Define the function

$$\delta(\eta) = egin{cases} -\eta + .35, & \eta < .35 \ 0, & \eta \geq .35 \end{cases}$$

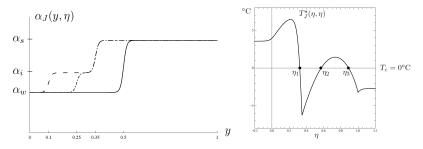
which represents the extent of the bare ice and the Jormungand albedo function

$$\alpha_{J}(y,\eta) = \frac{\alpha_{s} + \alpha_{w}}{2} + \frac{\alpha_{i} - \alpha_{w}}{2} \tanh(M(y-\eta)) + \frac{\alpha_{s} - \alpha_{i}}{2} \tanh(M(y-(\eta+\delta(\eta))))$$

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Jormungand albedo function and Temperature Profile



Widiasih's Theorem still applies with this albedo function
There is a locally attracting invariant manifold

$$\mathcal{P}_J^* = \{(\Phi_J^*(\eta), \eta) : \eta \in \mathbb{R}\}$$

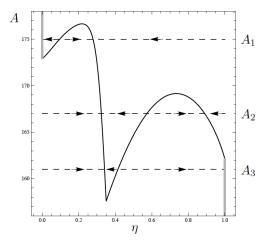
within $\mathcal{O}(\epsilon)$ of the manifold of fixed points

$$\mathcal{T}_J^* = \{(\mathcal{T}_J^*(y,\eta),\eta): \eta \in \mathbb{R}\}$$

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Jormungand Bifurcation in the Greenhouse Gas Parameter



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Insolation Distribution on Rapidly Spinning Planets

 Actual distribution can be found using orbital parameters (as seen in (5, 8, 17)):

$$s(y,\beta) = rac{2}{\pi^2} \int_0^{2\pi} \sqrt{1 - \left(\sqrt{1 - y^2} \sin\beta \sin\gamma - y \cos\beta\right)^2} \, d\gamma$$

• We use Legendre approximations in EBMs because the above integral doesn't have a closed form expression. Instead use

$$s(y,\beta) = \sum_{n=0}^{\infty} s_{2n}(\beta) p_{2n}(y)$$

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Insolation Distribution on Rapidly Spinning Planets

Write the 2n-th degree Legendre polynomial as

$$p_{2n}(y) = \sum_{k=0}^{n} a_{2k} y^{2k}$$

and

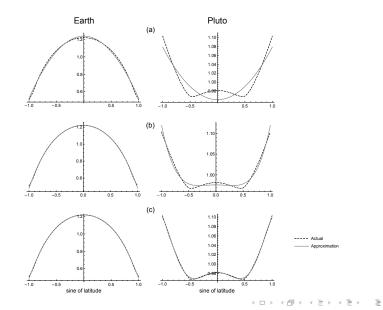
$$s_{2n}(\beta) = P_{2n} \sum_{k=0}^{n} a_{2k} c_{2k}(\beta).$$

where

$$c_{2k}(\beta) = \sum_{j=0}^{k} \binom{2k}{2j} \frac{(\sin\beta)^{2(k-j)}(\cos\beta)^{2j}}{\pi^{2}} \cdot \left(\int_{-\pi/2}^{\pi/2} (\cos\hat{\phi})^{2(k+1-j)}(\sin\hat{\phi})^{2j} d\hat{\phi} \right) \left(\int_{0}^{2\pi} (\cos\hat{\theta})^{2(k-j)} d\hat{\theta} \right)$$

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Insolation Distribution on Rapidly Spinning Planets



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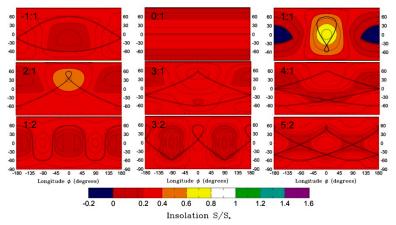
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Small Integer Spin-Orbit Resonances

Mean annual insolation distributions for obliquity $\beta = 60^{\circ}$ and eccentricity e = .2:



A. Dobrovolskis, "Insolation on exoplanets with eccentricity and obliquity."

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References

Open Questions about Insolation

 Can we quantify when we can use the "rapidly spinning planet" method/formula and still have error less than τ in our approximation?

• Can we find a closed form expression for insolation on planets with small integer resonances?

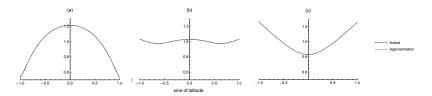
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Open Questions about Insolation

- For a reasonable range of parameter values, the Budyko map for Pluto doesn't have any nontrivial stable fixed points. Why?
 - Could Pluto's "upside down" insolation be playing a factor?
 - Is it because Pluto's insolation is relatively flat?



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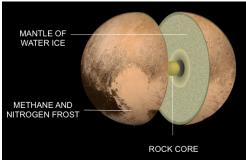
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Open Questions about Ice Lines

- Reformulate model to accommodate ice planets
 - Is more than one ice present and how do we account for different albedos?
 - How are the ices situated on the surface? Are ices mixed? Are they layered?



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Open Questions about Ice Lines

- Remove symmetry assumptions?
 - Investigate four ice lines $(\eta_{SP}, \eta_{SE}, \eta_{NE}, \eta_{NP})$ with the properties
 - (i) $\eta_{SP}, \eta_{SE}, \eta_{NE}, \eta_{NP} \in [-1, 1].$
 - (ii) $-1 \leq \eta_{SP} \leq \eta_{SE} \leq \eta_{NE} \leq \eta_{NP} \leq 1.$
 - (iii) Ice is located between η_{SP} and η_{SE} and between η_{NE} and η_{NP} .
 - Initial investigations into this case show potential oscillations in the ice line dynamics

Dynamics of the Ice Line

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Future Work

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References

Other Open Questions Concerning EBMs

- How do we accurately model a planet whose atmosphere "freezes out" (as Pluto's might)?
- Is the diffusion model more accurate for a planet with no oceans? For a planet without an atmosphere?
 - In which cases do we get the same results with both models?
 - In which cases do we get different results?

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(1)	M. Budyko. " 611-619.	The effect of solar radiation variation	ations on the climate of the Earth	." Tellus, 21 (5): 1969					
(2)		and G. North. "A Stability Theorem for Energy-Balance Climate Models." <i>Journal of the ic Sciences</i> , 36 : 1979. 1178-1188.							
(3)		and J. Coakley. "Analytical Analysis of a Budyko-Type Climate Model." <i>Journal of the Atmospheric</i> 32: 1975. 675-679.							
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